The Source and Fate of Massive Carbon Input During the Latest Paleocene Thermal Maximum

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Lithologic, faunal, seismic, and isotopic evidence from the Blake Nose (subtropical western North Atlantic) links a massive release of biogenic methane ~55.5 million years ago to a warming of deep-ocean and high-latitude surface waters, a large perturbation in the combined ocean-atmosphere carbon cycle (the largest of the past 90 million years), a mass extinction event in benthic faunas, and a radiation of mammalian orders. The deposition of a mud clast interval and seismic evidence for slope disturbance are associated with intermediate water warming, massive carbon input to the global exogenic carbon cycle, pelagic carbonate dissolution, a decrease in dissolved oxygen, and a benthic foraminiferal extinction event. These events provide evidence to confirm the gas hydrate dissociation hypothesis and identify the Blake Nose as a site of methane release.

Over a 10,000- to 20,000-year interval about 55.5 million years ago (1, 2), Earth’s climate and oceans warmed as deep-ocean and high-latitude surface water temperatures soared by 4° to 8°C (3–6), numerous mammalian orders (including primates) suddenly appeared in the geologic record (7), and many deep-sea species became extinct [foraminifera (3, 5, 8–12)] or disappeared temporarily [ostracodes (12)]. This extraordinary event, the latest Paleocene thermal maximum (LPTM) (4), coincided with a remarkable decrease in the ¹³C/¹²C ratio of global carbon reservoirs (2, 5, 8–11, 13–15) that has been explained by the addition of isotopically light CO₂ to the ocean, comparable to present-day fossil fuel input to the atmosphere (3, 15, 16). All of these observations have been linked by a single hypothesis (Fig. 1) involving a sudden change in ocean circulation and rapid escape of 1 × 10¹⁸ to 2 × 10¹⁸ g of CH₄ from marine gas hydrate reservoirs on continental slopes (11, 15–18). Nonetheless, without appropriate Paleocene continental slope sections, evidence for the theoretical CH₄ release has been lacking. Results presented here provide critical evidence to support rapid, massive CH₄ input to the ocean during the LPTM.

The “LPTM hydrate dissociation hypothesis” (15–18) can be summarized as follows (Fig. 1): Similar to the present day, vast quantities of CH₄ greatly enriched in ¹³C (³⁵³C ~ −60 per mil) were stored as gas hydrate in the upper few hundred meters of continental slope sediments (18–21) before the LPTM. Long-term global warming during the late Paleocene (22, 23) pushed the ocean-atmosphere system past a critical threshold (4, 17, 18), causing warm surface waters to sink and intermediate to deep ocean temperatures to rise by ~4° to 8°C (3, 8). This warming propagated into the sediments, converting once solid CH₄ hydrates into free gas bubbles (15). This warming resulted in an increase in pore pressure at depth, leading to sediment failure and the release of massive quantities of CH₄ into the ocean (19, 24, 25). Methane release would have occurred on continental slopes between 900 and 2000 m water depth because all gas hydrate in this sediment would be prone to dissociation in the Paleocene with a 4° to 8°C rise in bottom water temperature (Fig. 1) (15). The gaseous CH₄ reacted with dissolved O₂ (likely via bacterial activity (26)) to produce ¹²C-enriched CO₂, adding carbon to all reservoirs of the global exogenic carbon cycle and substantially shoaling the depth of carbonate dissolution in the ocean (15–18). Higher bottom water temperature, lower dissolved O₂, changes in surface water productivity, and more corrosive waters killed many of the deep-sea species (17). On land, higher partial pressure of CO₂ and elevated temperatures quickly opened high-latitude migration routes for the widespread dispersal of mammals (7). Over several hundred thousand years, global carbon and oxygen cycles gradually returned to equilibrium conditions after the LPTM (16, 18), although marine and terrestrial ecosystems were forever changed (7–12).

Ocean Drilling Program (ODP) Site 1051 on the Blake Nose (Fig. 2) recovered a continuous, expanded LPTM section (27, 28) that was deposited on the lower continental slope (present water depth, 1981 m). The unusually thick LPTM section yields sediment cycles that are caused by subtle lithologic variations and are best seen in magnetic susceptibility variations (Fig. 3). The cycles bundle into 100,000-year packages that each contain five shorter cycles, strongly suggesting a relationship to Earth’s precessional frequencies of ~20,000 years (27, 28). The precessional cycles provide an opportunity to study rates and timing of chemical and faunal changes associated with the LPTM, and to address the source and fate of CH₄.

At Site 1051, ~55% of benthic foraminiferal taxa disappeared in the latest Paleocene within a fraction of a precessional cycle (<5000 years) (Fig. 3). The majority (60%) of these last appearances occurs immediately below the carbon isotope excursion (CIE) and within a mud clast interval where faunal diversity plummeted from ~25 to 30 taxa to ~5 to 10 taxa per sample. Bulimina spp. dominate the surviving fauna, suggesting decreased dissolved O₂ conditions; at the same time, partly dissolved foraminiferal tests indicate increased ΣCO₂ and decreased CaCO₃ (Fig. 3). A similar rapid, pronounced benthic foraminiferal extinction event (BFEE) and environmental conditions mark the onset of the LPTM at other deep-sea locations (3, 5, 8–10) and probably resulted from a combination of factors, including increased temperature, decreased dissolved O₂, surface water productivity changes, and enhanced carbonate dissolution. Post-extinction reestablishment...
of benthic assemblages was gradual at Site 1051, with species origins occurring over a period of 200,000 years (Fig. 3).

Oxygen and carbon isotope records of a single foraminiferal group (Oridorsalis spp.) display large negative excursions across the LPTM at Site 1051 (Fig. 3), with δ18O and δ13C decreases of ~1.5 and ~3.0 per mil, respectively, across a ~20-cm interval. Bulk sediment δ13C values decrease by ~1.2 per mil across this interval (Fig. 3), although we did not measure bulk samples from the interval of peak δ13C excursion. At Site 1051, the onset of the isotopic excursions coincides with the BFEE and is contemporaneous with similar isotopic changes in other deep-sea LPTM sections from around the world (3, 5, 17).

The benthic foraminiferal δ18O decrease at Site 1051 indicates an increase of >6°C in bottom water temperature on the continental slope (29). The onset of the δ18O change occurs over 20 cm (or one-quarter to one-third of a single precessional cycle; Fig. 3), indicating that substantial warming of intermediate water occurred within 5000 to 7000 years during the LPTM. The amount, duration, and location of bottom water warming inferred at Site 1051 is sufficient to dissociate large quantities of gas hydrate in middle to lower slope sediments (15, 30).

The benthic foraminiferal CIE at Site 1051 is similar in shape and magnitude to those in other LPTM deep-ocean sections (3, 5, 8, 17) and represents a decrease of 3 per mil in the δ13C of ΣCO2 in all water beneath the thermocline. Major decreases in δ13C records of planktonic foraminifera or bulk sediment carbonate (dominated by tests of surface-dwelling organisms) at Site 1051 and other marine locations (3, 5, 10, 14) indicate a large decrease in the δ13C of surface water ΣCO2. A decrease of 4 per mil in the δ13C of terrestrial carbonate and organic carbon also has been documented (11, 13). Thus, the prominent CIE across the LPTM occurred in shallow marine, deep marine, and terrestrial reservoirs, indicating massive carbon addition to the global exogenic carbon cycle (5, 14–16), as opposed to carbon redistribution within this system (31–33). Pronounced carbonate dissolution on the sea floor (5, 9–11, 14, 17), including at Site 1051 (Fig. 3), is consistent with this interpretation (15–18). The duration of carbon input is expected to be equivalent to the onset time of the CIE (18), or about 5000 to 7000 years according to the δ13C records at Site 1051 (Fig. 3). Release and oxidation of 1 × 1018 to 2 × 1018 g of CH4 and the subsequent propagation of CO2 through various carbon reservoirs, is the only known mechanism to explain the sudden, extreme, and global nature of the CIE (11, 15).

Model results (18) show that the release and oxidation of 1.2 × 1018 g of CH4 over 7500 years in the present-day deep Atlantic would result in a rapid decrease of the δ13C of ΣCO2 in the deep Atlantic, followed by an exponential recovery over a period of about 140,000 years as excess carbon is eventually transferred to the rock cycle (Fig. 4). The temporal resolution afforded by precessional cyclicity in the expanded section at Site 1051

Fig. 2. Multichannel seismic line TD-5. (A) Map showing location of inset (B). BBOR, Blake Bahama Outer Ridge. (B) Enlarged location map with drill sites and seismic lines shown in (C) and (D). (C) Chaotic reflections correlated with the CIE near an Aptian reef may be a CH4 release site. (D) Erosional scarp and current-controlled sedimentation associated with a reflection correlated with the CIE; reflection A is the equivalent age to reflection A mapped on the western Bermuda Rise (37). Modified from (27).

Fig. 3. Benthic foraminiferal extinctions (HO, highest occurrence) and decrease in diversity (number of taxa per sample) are associated with the CIE and deep-water warming (indicated by benthic isotopes, Oridorsalis spp.), whereas origins (LO, lowest occurrence) are scattered throughout. Triangles were derived from a single large specimen and are considered outliers. Magnetic susceptibility (related to ~20,000-year precessional cycles) provides a chronology. Mud clasts indicate sediment failure that may have resulted from CH4 release; CH4 oxidation led to the overlying dissolution interval and low O2 benthic fauna (mbsf, meters below sea floor).
low waters; instead, it contains partially compacted, fine-grained clasts up to 5 cm long and has a scour mark at its base (27). The lithologic, faunal, and geochemical evidence indicates that the intraformational mud clast layer was transported from a penecontemporaneous horizon just upslope from Site 1051 at the crucial time, precisely at the onset of the CIE (Fig. 3). The juxtaposition of the top of a mud clast layer and the base of the CIE in a sediment succession deposited on the lower slope is a compelling argument for CH$_4$ release via sediment failure at the Blake Nose during the LPTM. Evidence for the requisite deep-water warming just before the CH$_4$ release is not evident in our data, but would have been erased by erosion that formed the scour mark at the base of the mud clast interval.

Seismic profiles across the Blake Nose (27) reveal three features (Fig. 2) associated with the CIE: an extensive interval of chaotic reflections above a buried Aptian reef, a prominent seismic reflection, and a scarp upslope from Site 1051. Buried reefs may serve as traps for overpressured gas and gas hydrates (36), and CH$_4$ could have vented from the region of chaotic reflections during the LPTM (Fig. 2). This area was at the appropriate depth given for slope failure caused by hydrate dissociation in the Paleocene (Fig. 1). The chaotic interval is associated with a seismic reflection traced to Site 1051 that correlates with the CIE. This seismic reflection apparently is caused by the impedance contrast between the mud clast layer and the overlying dissolution interval observed at Site 1051. The CIE reflection clearly postdates reflection $A'$, a time when widespread current-controlled sedimentation began in the western North Atlantic (37). A scarp upslope from Site 1051 (Fig. 2) may be related to slope failure associated with CH$_4$ release or to erosion associated with current-controlled sedimentation. We cannot discriminate between the two possibilities with seismic profiles available at present, but we note that the chaotic zone, seismic reflection, and scarp can be traced to the depth of the CIE at Site 1051 (Fig. 2), indicating a temporal equivalence and suggesting a causal link.

Even though our results suggest that CH$_4$ was released from the Blake Nose region during the LPTM, the mass of CH$_4$ from this region alone is insufficient to explain the magnitude of global perturbations at the LPTM. Other sections deposited on the middle to lower slope during the LPTM also must exhibit features similar to those reported here.

References and Notes
27. R. D. Norris et al., Proceedings of the Ocean Drilling Program, Initial Reports, 1778 (College Station, TX, 1998).

16 July 1999; accepted 19 October 1999

Fig. 4. Comparison of theoretical (18) and observed (Site 1051) LPTM carbon isotope changes.