

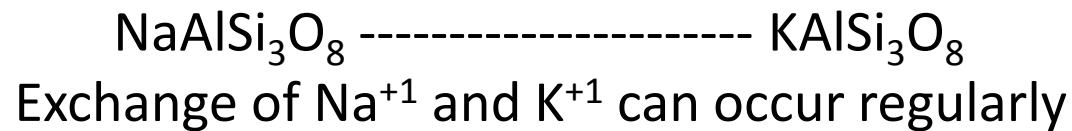
**Illustrating mineral chemistry:
graphical displays of
multicomponent systems**

Initial concepts: two freely substituting components

End-members: two idealized starting components (cations) that could ideally fill the same site in a mineral structure

This leaves the possibility that a mixture of these components could occur

For example, the alkali-feldspars



Plotting the ratio of the cations can be done independently or the relationship can be plotted based on the relative wt% of each species

3 ways of expressing mineral composition

Talc (sheet silicate usually formed from metamorphosed ultramafics)

In terms of weight percent oxides (wt%)

I1

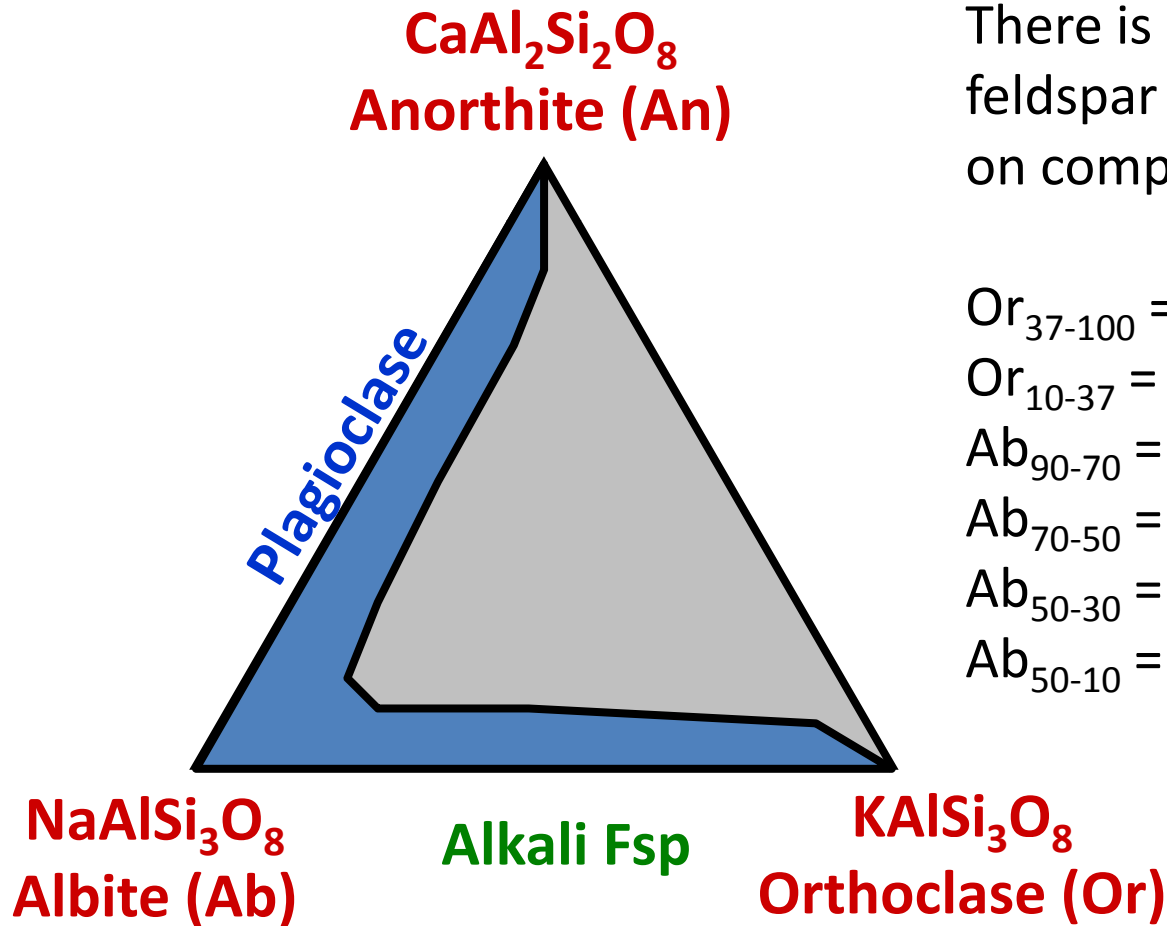
Chemical formula	MgO	SiO ₂	H ₂ O	MgO	SiO ₂	H ₂ O
Mg ₃ [Si ₄ O ₁₀](OH) ₂	31.88	63.37	4.75	3	4	1

In terms of wt% elements and apfu'

Mg	Si	O	H	Mg	Si	O	H
19.23	29.62	50.62	0.53	3	4	12	2

Feldspar Group

Feldspars are the most common silicates in the Earth's Crust.



There is an “old” system for feldspar nomenclature, based on composition.

Or₃₇₋₁₀₀ = Sanidine

Or₁₀₋₃₇ = Anorthoclase

Ab₉₀₋₇₀ = Oligoclase

Ab₇₀₋₅₀ = Andesine

Ab₅₀₋₃₀ = Labradorite

Ab₅₀₋₁₀ = Bytownite

The Ternary diagram

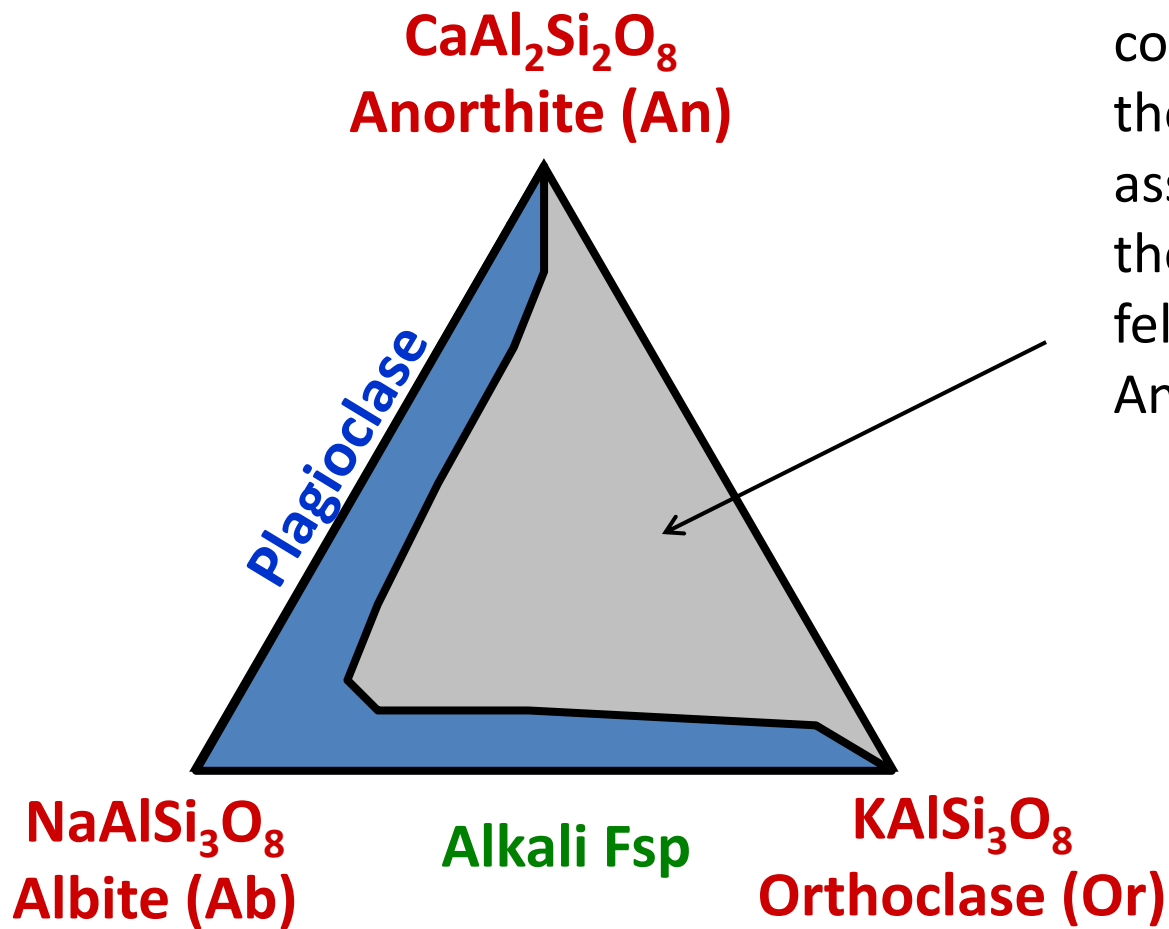
Three distinct end-members: each corner of the diagram is 100% of one component and 0% of the other two

Usually ternary plots are divided into a triangular gridded set with each line representing a 10% composition change

Returning to the feldspars:

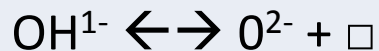
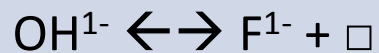
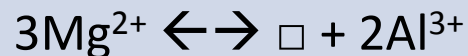
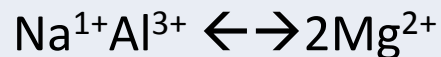
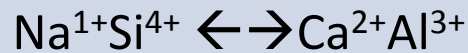
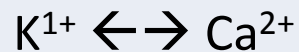
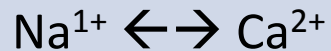
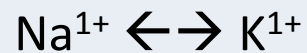
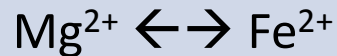
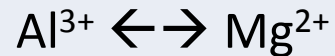
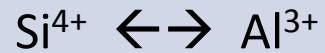


In this case, the ternary diagram represents two distinct but inherently linked solid-solution series



Miscibility gap: range of composition(s) where there are no stable assemblages in nature. In the example of the feldspars, there is no Or – An solid-solution series

COMMON SUBSTITUTIONS



One for one cation substitution are a very common manner by which chemical formulas can be altered; however, notice the charge balance is often not maintained.

In these cases, additional cation substitutions and/or a change in the number of anions present in the formula is necessary to re-achieve an equilibrium electrical state.

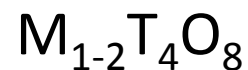
Structural sites in minerals

Based on the size and charge of cations, we already know that there are particular types of coordination polyhedra based on the size of the cation(s) that must be accommodated, there are specific locations in a mineral where particular species will and will not fit.

Just as with creating or individual coordination polyhedra, there are specific sites in a mineral assemblage that are the appropriate size and capacity for a specific set of elements

These locations are given notations so that we can write generalized mineral formulas for a related mineral family irrespective of the specific mineral composition

Returning again to the feldspars:



Structural sites in minerals

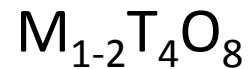
Most simply the terms we will use for the major sites are:

M (1-4): “metal” cation site

T: tetrahedral site

A, B, C: interstitial sites between coordinated locations

Returning again to the feldspars:



Generalize formulas and sites for the major minerals

Garnet:



A > B and

A = Ca²⁺, Mg²⁺, Fe²⁺, Mn²⁺

B = Al³⁺, Fe³⁺, Cr³⁺



Generalize formulas and sites for the major minerals

Olivine:



$M_2 > M_1$ and

$M_2 = Ca^{2+}, Mg^{2+}, Fe^{2+/3+}, Mn^{2+}$

$M_1 = Mg^{2+}, Fe^{2+/3+}, Mn^{2+}$

fayalite (Fe_2SiO_4) vs. monticellite ($CaFeSiO_4$)



An example from the Olivines

Olivine is a solid solution series between an Mg and Fe rich set of end-members: Forsterite and Fayalite

The general formula for Olivine is: $(\text{Mg, Fe})_2[\text{SiO}]_4$

Chemical analysis (from Floran & Papike, 1973)

SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	FeO	MnO	MgO	CaO
30.09	0	0	0	69.42	0.28	0.91	0.08

Which Olivine is this an analysis of?

How can we calculate the formula in apfu'?

How many oxygen atoms will we need to base the calculation on?

An example from the Olivines

Olivine is a solid solution series between an Mg and Fe rich set of end-members: Forsterite and Fayalite

The general formula for Olivine is: $(\text{Mg, Fe})_2[\text{SiO}]_4$

Based on the general formula, we can now assign a new formula to this specific species of Olivine:



Sum of M1 and M2 cations: $0.0453+1.937+0.008+0.003 = 1.9933$ (~2)

Octahedral coordination polyhedra of the mineral [fayalite](#)

Generalize formulas and sites for the major minerals

Pyroxene:



M2 > M1 and

M2 = Ca^{2+} , Na^{1+} , Fe^{2+} , Mg^{2+}

M1 = Mg^{2+} , $Fe^{2+/3+}$, Mn^{2+} , Al^{3+} , Cr^{3+} , Ti^{4+}

enstatite ($Mg_2Si_2O_6$) vs. diopside ($CaMgSi_2O_6$)



Generalize formulas and sites for the major minerals

Amphibole:



$A > M4 > M3 \sim M2 = M1$ **or** $A > B > C > T$ **and**

$A = K^{1+}, Na^{1+}$

$B = Na^{1+}, Ca^{2+}, Mg^{2+}, Fe^{2+}, Mn^{2+}$

$C = Mg^{2+}, Fe^{2+/3+}, Mn^{2+}, Al^{3+}, Cr^{3+}$

$T = Si^{4+}, Al^{3+}, Ti^{4+}$

$OH = (OH)^{1-}, F^{1-}, Cl^{1-}, O^{2-}$

Anthrophyllite ($Mg_7Si_8O_{22}(OH)_2$) vs. Tremolite ($Ca_2Mg_5Si_8O_{22}(OH)_2$) vs. Richterite ($Na(CaNa)Mg_5Si_8O_{22}(OH)_2$)