



ABOVE: The major watersheds of Lake Tuborg (~81°N 75°W). The southwest basin is meromictic (about 155 m deep): the northeast basin is entirely freshwater (about 80 m deep). Watershed A is nivally fed, B is mostly nivally fed, C and D are mostly glacially fed. Inset shows location of Lake Tuborg on Ellesmere Island.



ABOVE: (A) 1959 aerial photograph of Lake Tuborg highlighting Delta A. (B) Bathymetric map of the study area near Delta A. Black symbols are points where bathymetry was measured. Red symbols are CTD cast sites. The red grid illustrates the sampling area. White labels are depth (m) below lake level.

INTRODUCTION: Lake Tuborg is a large glacially fed and glacially dammed lake on west-central Ellesmere Island. It has two sub-basins. The southwest sub-basin is meromictic, and has a maximum depth of 155 m. The northeast sub-basin is freshwater and is shallower (Zmax, ~80 m).

This poster shows processes of sedimentation recorded near a large nival delta (watershed A at left, photo at right) located at the north-central portion of the lake. Processes were recorded during peak nival melt in 2001 with sediment traps, a datalogging flow meter, and a CTD.

Annual suspended sediment discharge from this nival delta is likely much less than glacially-fed deltas at Lake Tuborg. However, the nival tributary flows into the lake in its meromictic basin, and the delta face is steep. Therefore, this is an ideal location to study how sedimentary processes are affected by meromixis.





LACUSTRINE SEDIMENTARY PROCESSES NEAR A HIGH ARCTIC DELTA, LAKE TUBORG, ELLESMERE ISLAND Ted Lewis (lewist@geo.umass.edu); Raymond. S. Bradley; Pierre Francus University of Massachusetts, Amherst. Climate System Research Center



(C) Current meter courtesy of Roger Lewis and John Sweeney. Based on CTD profiles, the difference in concentration between reshwater and saltwater based on salinity and temperature is about 20 g/L. During the 1995 slush flow at Delta D, SSC reached 5 g/L (Braun, 1997), while SSC reached 2 g/L during the 2001 nival flood at Delta A. This theoretically precludes the occurrence of fluvially generated underflows in the meromictic sub-basin of Lake Tuborg, even during extreme events.



Above: Weather and climate at Lake Tuborg and regional weather stations.

(A) Precipitation at Eureka, 1948-2001. Lower graph (black) shows mean precipitation for each day. Upper graph (red) shows the number of days used for to calculate mean precipitation.

(B) Red line graph shows mean daily 2001 air temperature recorded at lake level with a datalogging thermistor. Temperature was recorded with a 15-minute interval. Black line graph shows 1947-2001 air temperatures recorded at Eureka. Black bar chart shows 2001 Lake Tuborg precipitaton (mm).

(C) Lake level fluctuations at Lake Tuborg in 2001. Zero datum is the lowest recorded lake level.

(D) Air temperatures recorded at Lake Tuborg, Eureka and Resolute in 2001. R-squared for Resolute and Tuborg is 0.57; R-squared for Eureka and Tuborg is 0.90.

Mean daily temperature (5/21-8/11) for Lake Tuborg , 2.9 *C Resolute. 2.9 *C



Rain Gauges



Radiation Shield

Range











with a Sea Bird Instruments SBE19 CTD (Conductivity Temperature/Transmissivity Depth) profiler with a 660 nm. 25 cm path length SeaTech transmissometer

	Temperature	Conductivity	Pressure	Transmissivity
	-5 to +35 °C	0-7 S/m	50 to 10,000 psia	0-5 V (~0.004-20 mg/L)
	0.01 °C /6 months	0.001 S/m/month	0.5% of full scale	?
1	0.001 °C	0.0001 S/m	0.03% of full scale	0.001 V





Overflow plumes

nputs.

105 m apart

(attenuation increases)

adjacent to tributary

At left: (A) Attenuation on June 16, 2001. Note the large attenuation increase at the chemocline (~55 m) This attenuation spike is seen throughout the meromictic basin even before the onset of summer snowmelt. This could be a result of algal mats, sediment of insufficient density to pass into the monimolimnion, or precipitation from the dissolved load when anoxic water (in the monimolimnion) comes in contact with oxic water (in the mixolimnio (B) Attenuation on June 28. The attenuation spike i thicker. less well defined, and has shallowed somewhat. Meltwater may have caused some mixing, and suspended sediment may have "ramped

on the density difference at the chemocline. (C) Water temperature on June 19. Note the relatively warm water of the monimolimnion and mixing near the delta face. (D) Attenuation on June 19. Turbid water near the



At left: (E) Attenuation on June 19. (F) Water temperature on June 19. The overflow plume is composed of water slightly warmer than its surroundings,

At right: Top: Each 3-D visualization is composed of volumetric pixels (voxels) Each voxel has an interpolated value. Histograms of attenuation for each date were produced. **Bottom:** A surface plot combining the four histograms (cf. Beierle et al., 2002).

Each voxel has a known volume, and field calibration of the transmissometer to suspended sediment concentration should be possible. Therefore, the mass of sediment within each visualization (or a portion of each visualization) could be calculated



See http://www.geo.umass.edu/gradstud/lewist/lewist.htm and http://www.paleoclimate.org for more details.



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